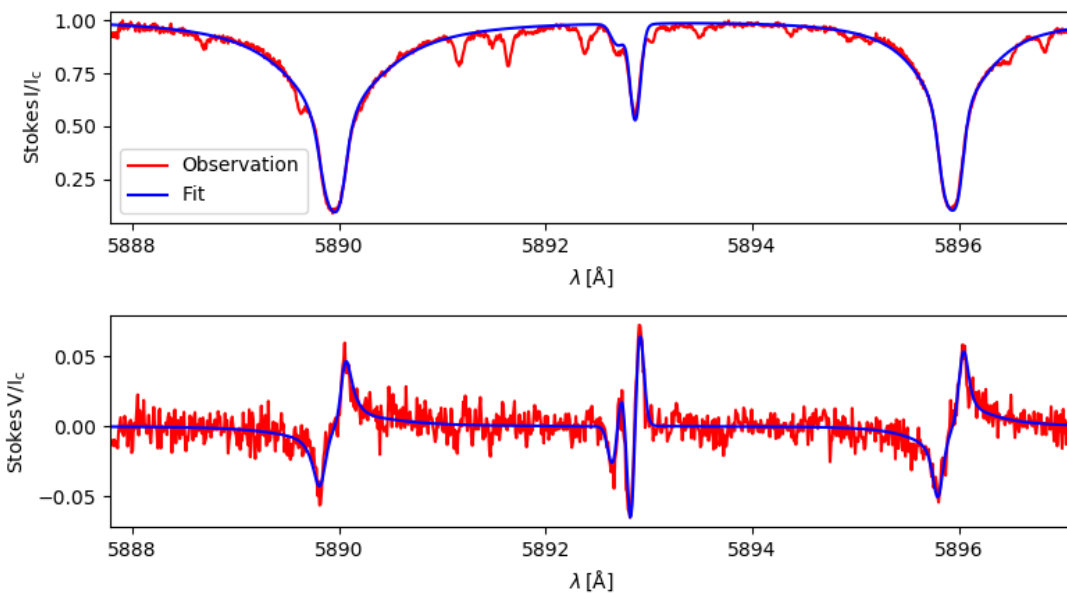


PHYS 7810: Solar Physics with DKIST

Lecture 21: Inversions and how to do them

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Previous lectures

- We have been introduced to line formation (model atmosphere → spectrum)
- We have been introduced to model fitting
- Once you choose your **model** and have your **observables** you can perform inference
- We have seen that method of choice is **posterior maximization** that with uniform priors reduces to **maximum likelihood**, which under the assumption of gaussian errors reduces to **chi-squared minimization**.

$$\chi^2 = \sum_i \sum_s (I_{i,s}^{\text{obs}} - I_{i,s}^{\text{synth}})^2 / \sigma_{i,s}^2$$

What we usually omit is what synthetic intensity depends on

- Our model is our choice of the approach to model polarized line formation.

$$I_{s,i}^{\text{synth}} = \mathcal{F} \left[T(h), p_g(h), v_{los}(h), \zeta(h), \vec{B}(h) \right]$$

Some approaches

(increasing complexity, let's discuss each one separately)

- Weak field approximation (what we did last class)
- Milne-Eddington approximation (Everything constant, the source function is linear)
- ***Stratified atmosphere, with all species in LTE***
- ***Stratified atmosphere, with species in NLTE***
- Stratified atmosphere, NLTE, with Partial Frequency Redistribution
- “Non atmospheric” models, e.g. slab model for hazel

We are going to focus on middle two

- ***Stratified atmosphere, with all species in LTE***
- ***Stratified atmosphere, with species in NLTE***
- We have observed spectra, we want to know atmospheric structure.
- But that is a function.
- Simplest is to discretize somehow. Then we can solve forward problem using methods from previous classes.

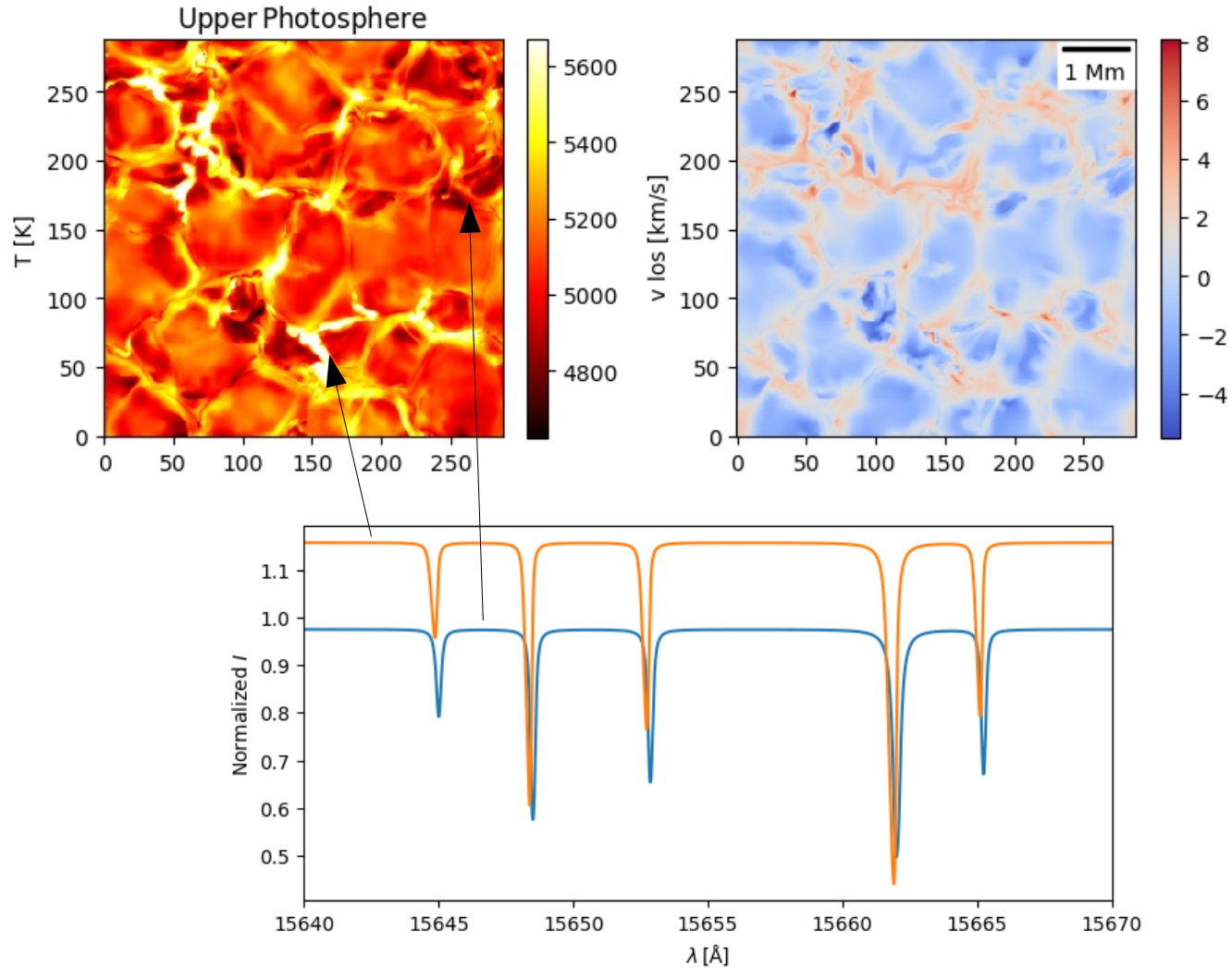
LTE line synthesis flowchart:

Atmosphere model: temperature, density, velocity, magnetic field

Use Saha-Boltzmann to calculate all relevant level populations

Calculate opacity/emissivity everywhere in the atmosphere

Solve Radiative Transfer Equation → Get the spectrum



What now, what comes next?

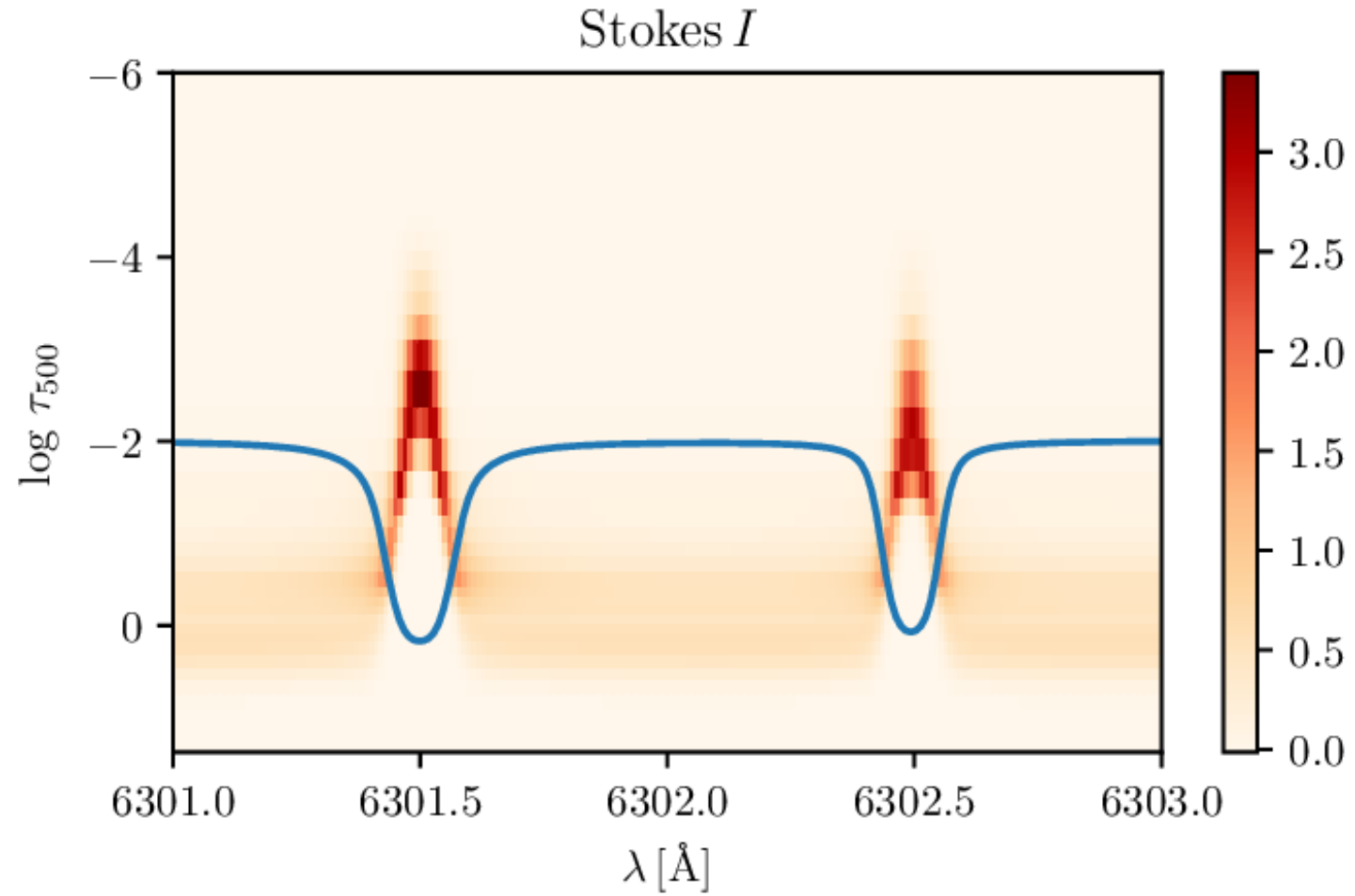
- We have to change the parameters \mathbf{p} of the atmosphere to get a fit

For us it would be:

$$\sum_j^{NM} \frac{\partial I_{s,i}(\mathbf{p})}{\partial p_j} = (I_{s,i}^{\text{obs}} - I_{s,i}^{\text{synth}}) = \delta I_{s,i}$$

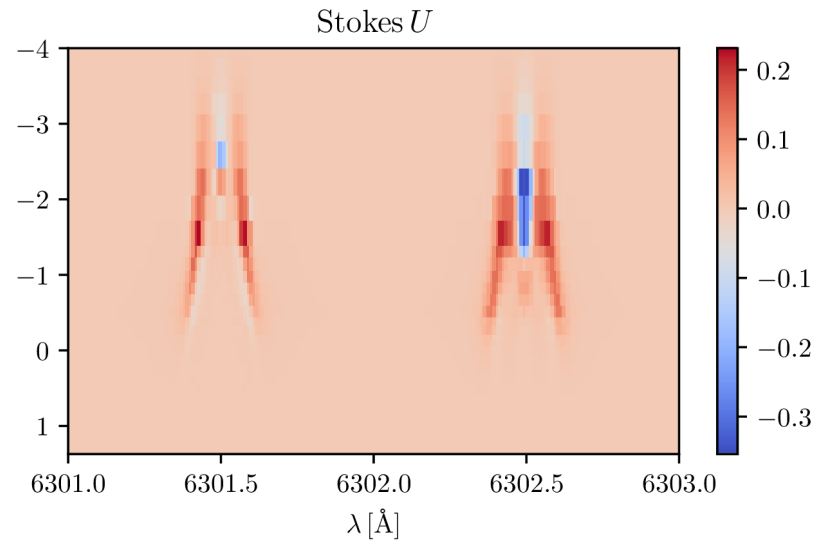
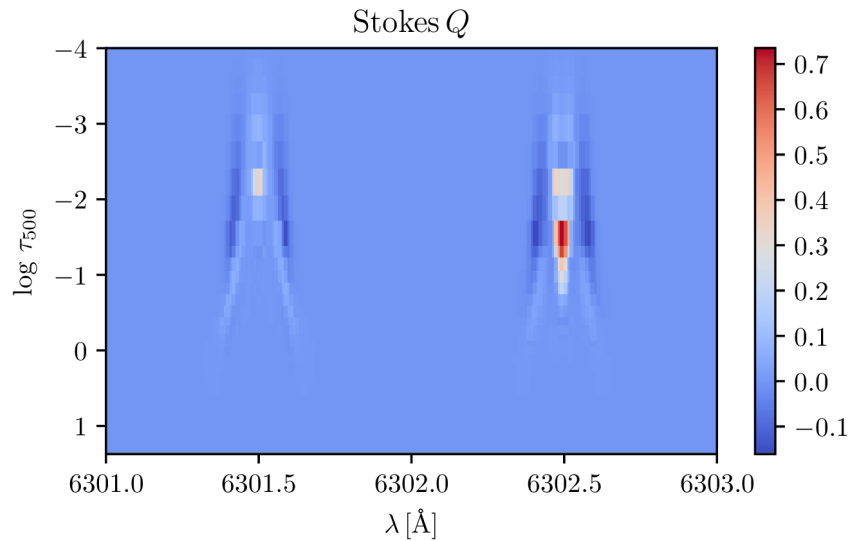
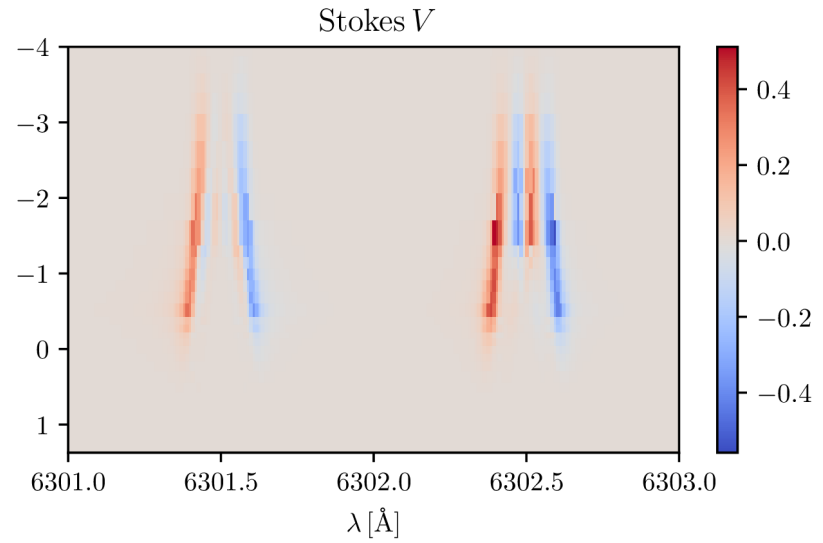
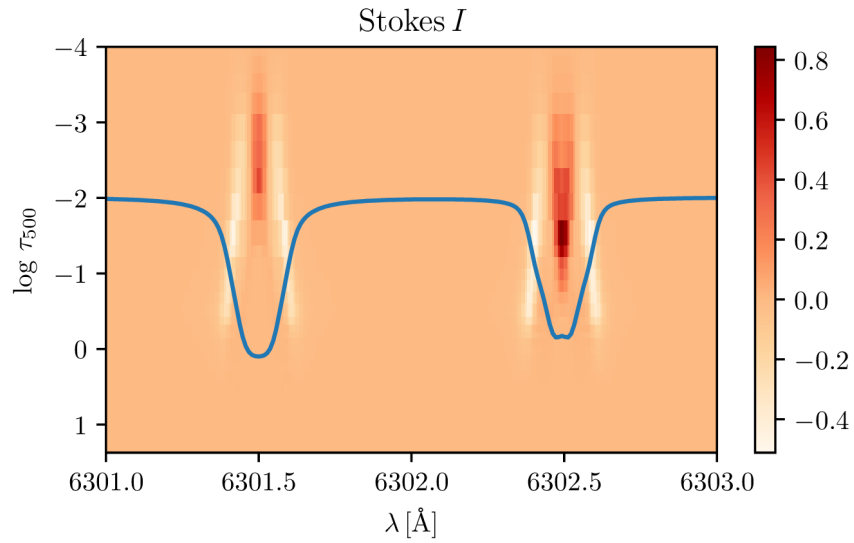
We will call the partial derivative of the specific polarized intensity with respect to a physical quantity at given depth **response function**.

Again, the response function, let's talk a lot about these



$$\frac{\partial I(\lambda)}{\partial T_d}$$

They look different for different parameters, eg. B



Levenberg – Marquardt

$$\sum_j^{NM} \frac{\partial l_{s,i}(\mathbf{p})}{\partial p_j} = J_{j,i} \delta p_j = \delta l_i$$

(I absorbed s into i to be more concise, as we did few slides back).

Scheme:

- Assume a model atmosphere, i.e. vector \mathbf{p}_0 .
- Calculate $\mathbf{l}(\mathbf{p})$ and evaluate $\hat{\mathbf{J}}$ and $\delta \mathbf{l}$.
- Calculate correction to the parameter vector as by multiplying both sides with $\hat{\mathbf{J}}^T$:

$$\hat{\mathbf{J}}^T \hat{\mathbf{J}} \delta \mathbf{p} = \hat{\mathbf{J}}^T \delta \mathbf{l}$$

- Correct the model and re-start the loop. As soon as \mathbf{l}^{calc} is good enough, stop.

How to calculate the response functions?

We imagine we perturbed a parameter at some depth. Everything changes.
What we want to know is how intensity changes

$$\frac{dI}{dz} = -\chi I + j$$

$$\frac{dI + \delta I}{dz} = -(\chi + \delta\chi)(I + \delta I) + (j + \delta j)$$

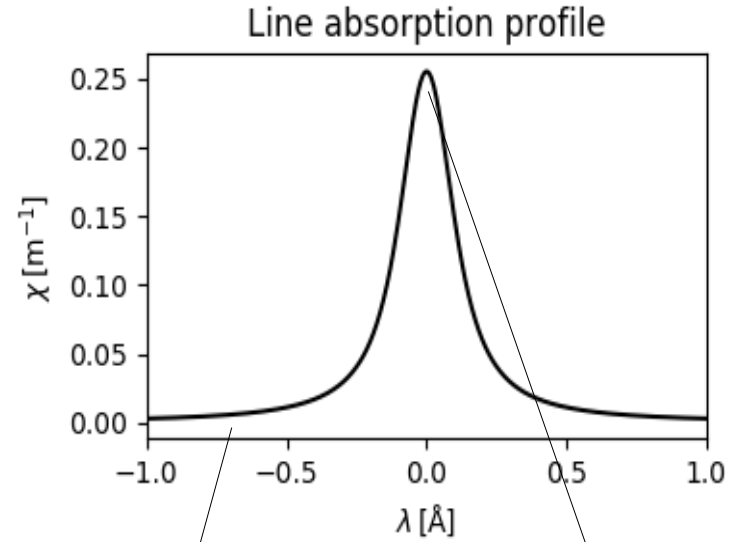
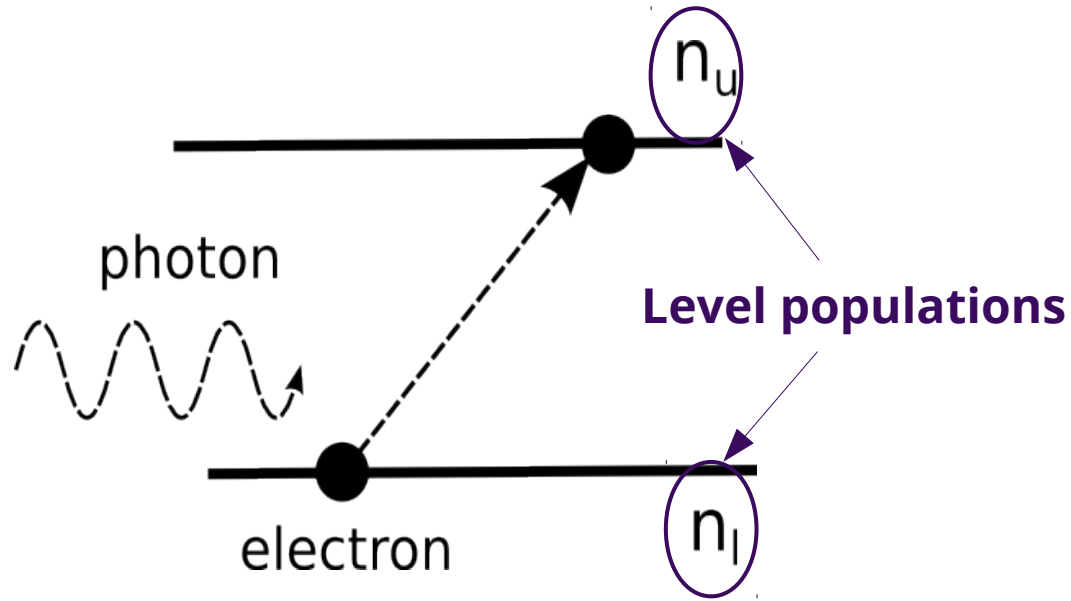
Change
(response) in
the intensity
(what we need)

$$\frac{\delta I}{dz} = -\chi\delta I + (\delta j - \delta\chi I)$$

This looks like RTE!

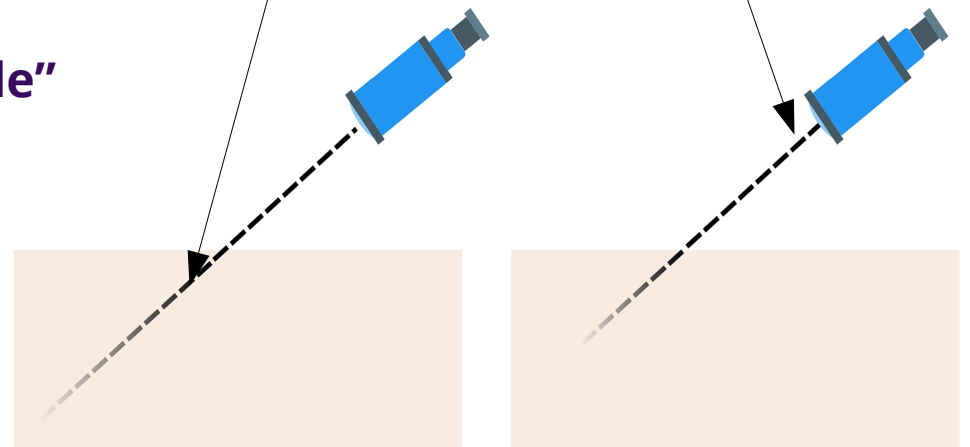
Changes in
opacity and
emissivity (what
we have to
move on to)

Opacity and emissivity responses follow from here



$$j(\lambda) = \frac{hc}{4\pi\lambda} n_u A_{ul} \phi(\lambda)$$
$$\chi(\lambda) = \frac{hc}{4\pi\lambda} (n_l B_{lu} - n_u B_{ul}) \phi(\lambda)$$

The term $\phi(\lambda)$ in both equations is circled and labeled "Line profile".



So, we need “responses” of opacity and emissivity

- There are two parts. **“Obvious one”**, changes in the line profile (splitting, shifting, broadening, polarization)
- **Non-obvious one: changes in the level populations (extremely complicated in NLTE)**

$$\frac{n_{j+1}n_e}{n_j} = \frac{2Z_{j+1}}{Z_j\Lambda^3} e^{-E_{\text{ion}}/kT}$$

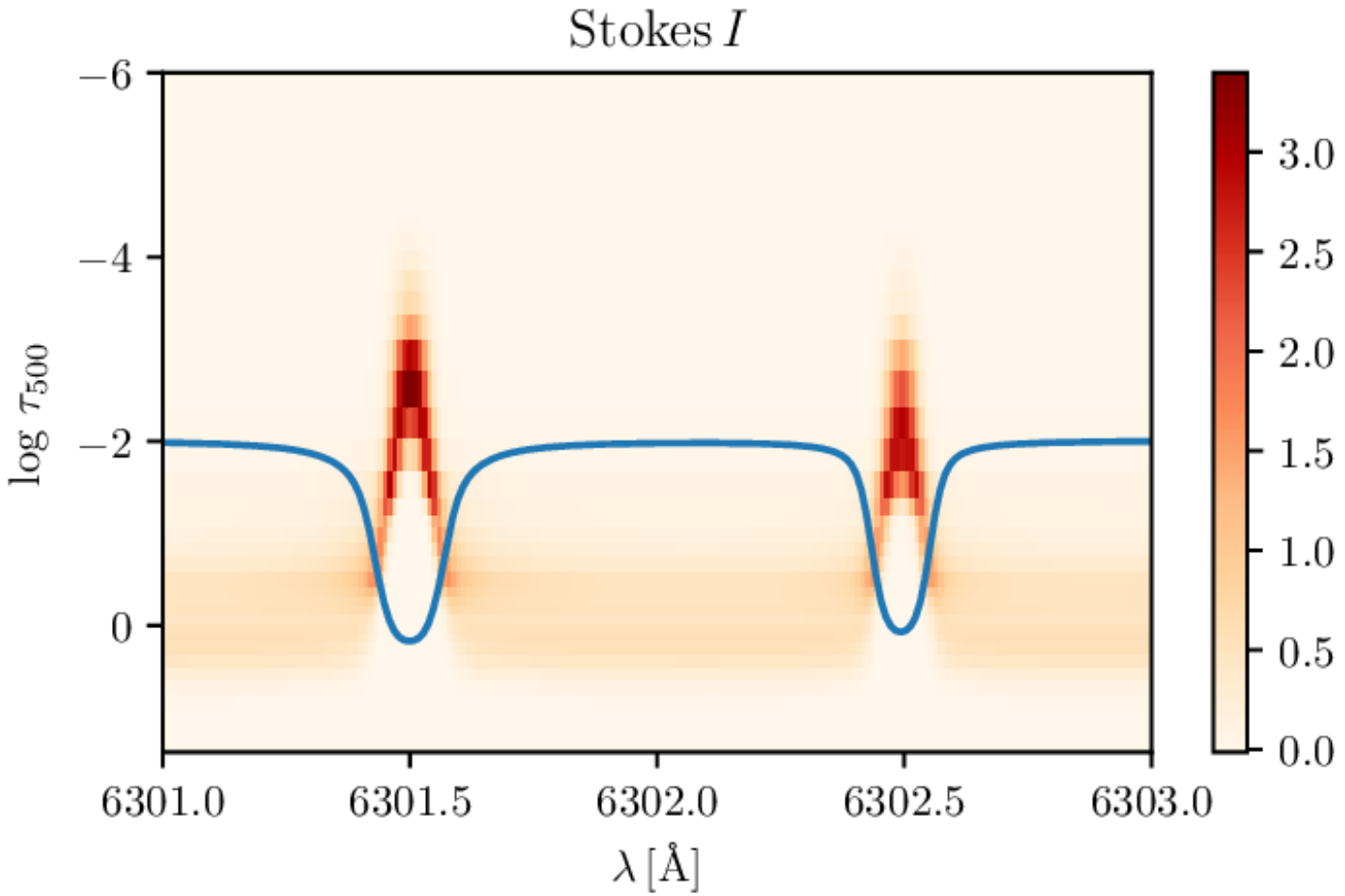
$$\Lambda = \sqrt{h^2/2\pi m_e kT}$$

$$n_{j,i} = n_j \frac{g_i e^{-E_i/kT}}{Z_j}$$

Response function calculation recipe

- Using Saha equation, calculate how the ionization changes and molecular number densities change when we change parameters (temperature, density)
- Using Boltzmann equation, see how excitation changes (temperature)
- Using your choice for damping see how damping changes (temperature, density)
- See how the line shift (los velocity)
- And broaden (temperature, microturbulence)
- And split / polarize (magnetic field vector)
- Put them all together, and “propagate” to get response function.

You need to do this once for each depth point to get this

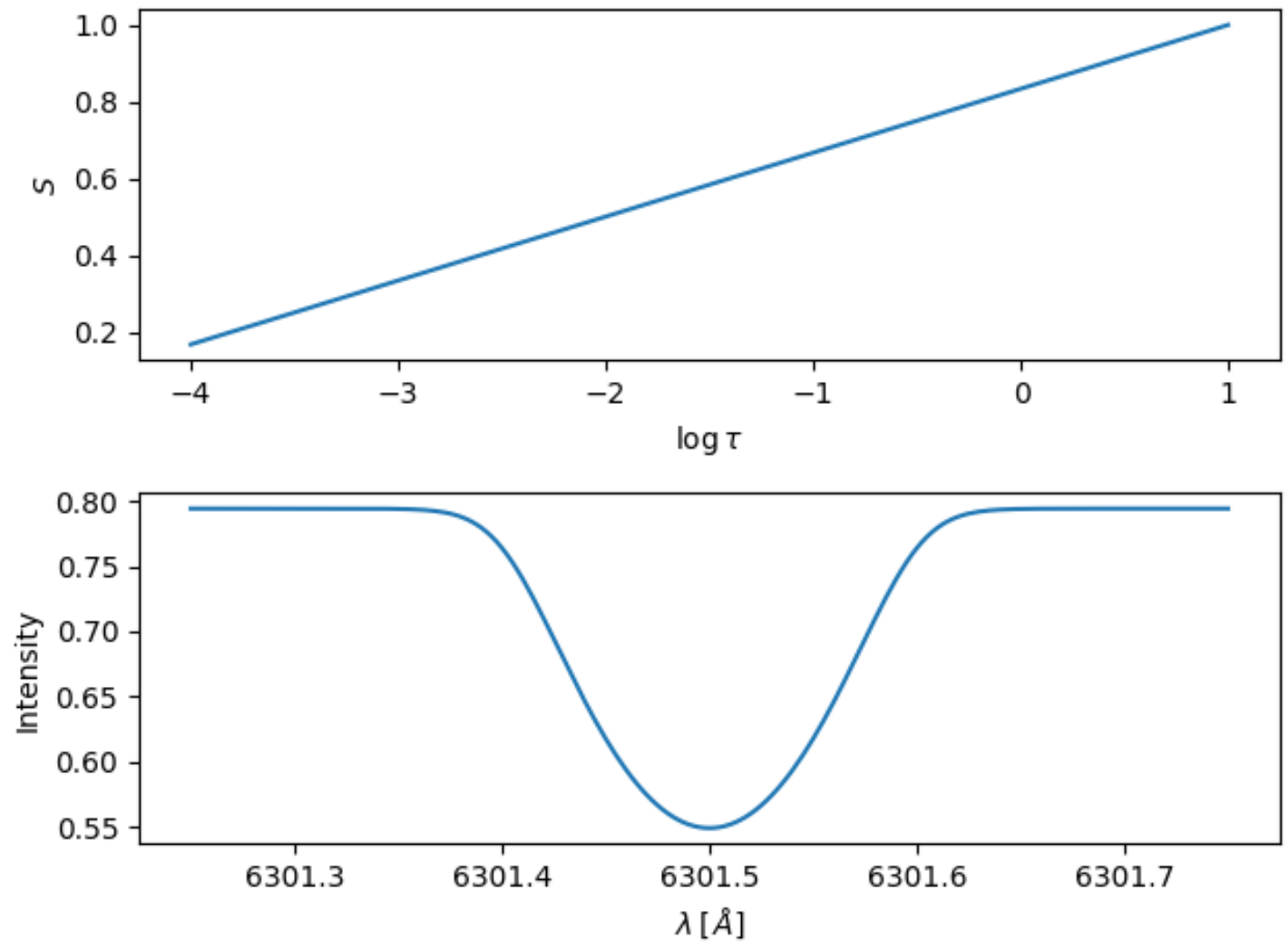


$$\frac{\partial I(\lambda)}{\partial T_d}$$

But now we just plug this into the L-M and we are done, right?



A very simple linear line formation model to illustrate this



My model is:

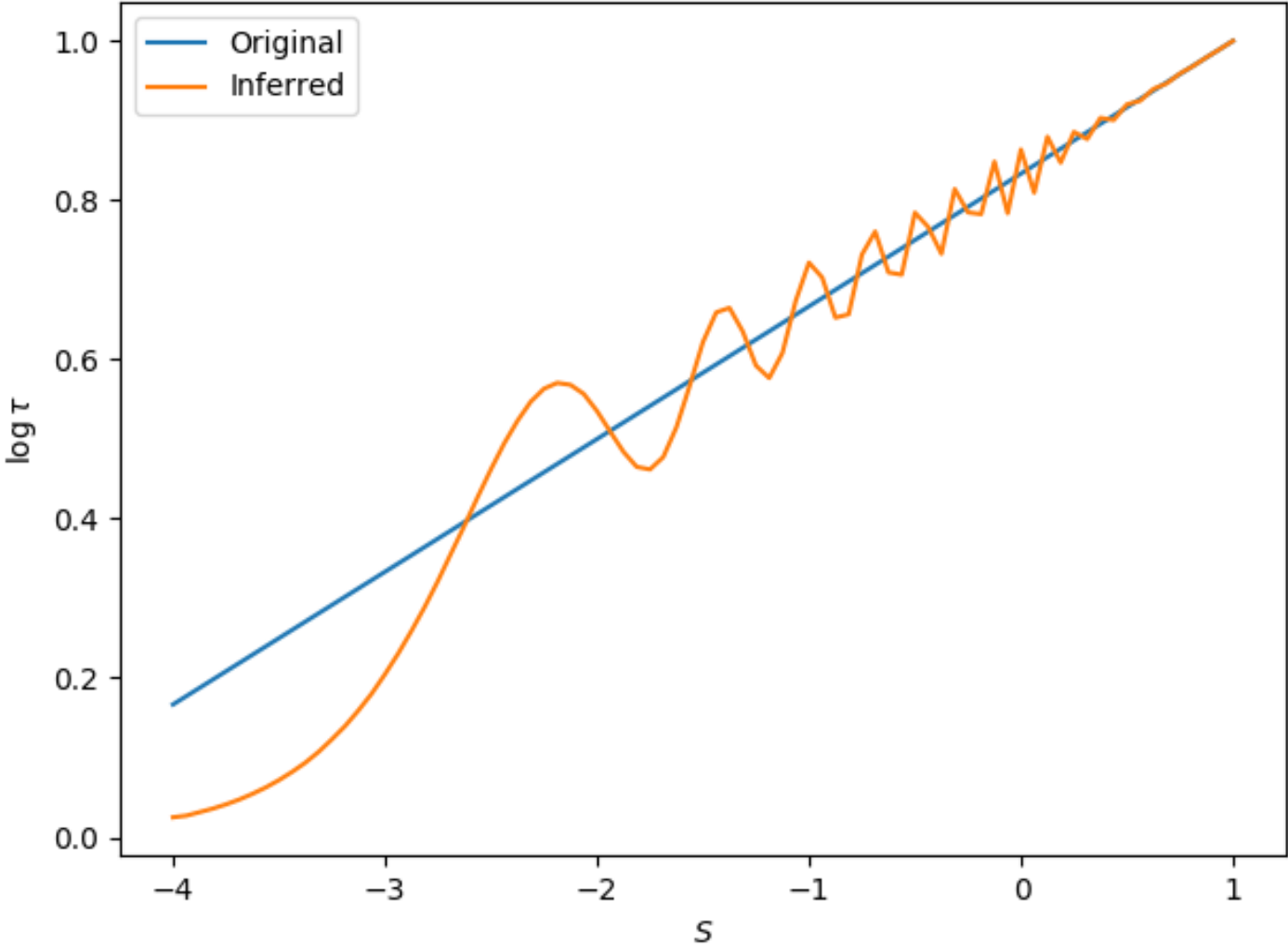
Measurement

Unknown

$$I_i = \sum_d w_{d,i} S_d$$

Weights, we assume we know (we don't in general!)

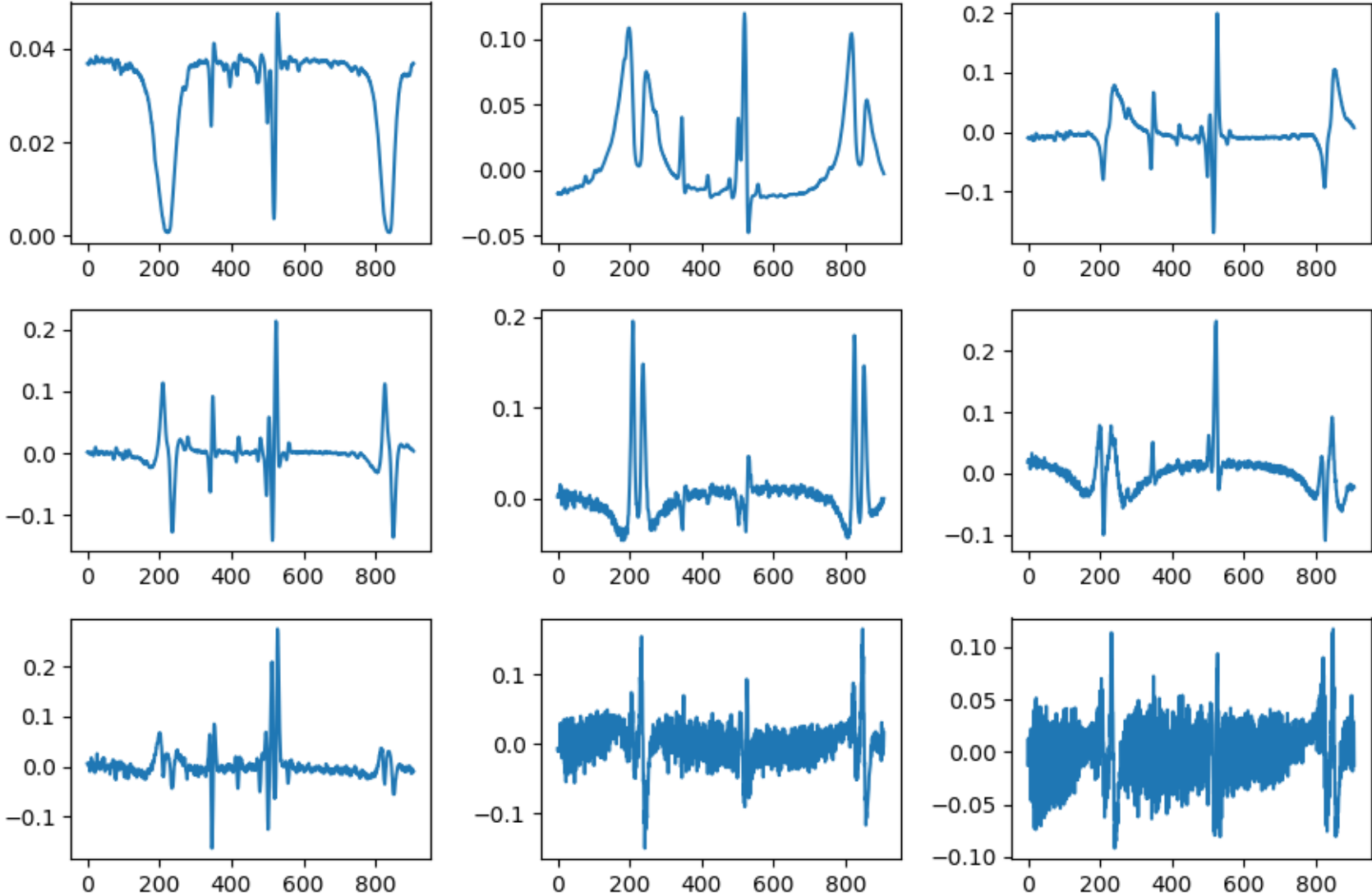
Does not work



Why is this so?

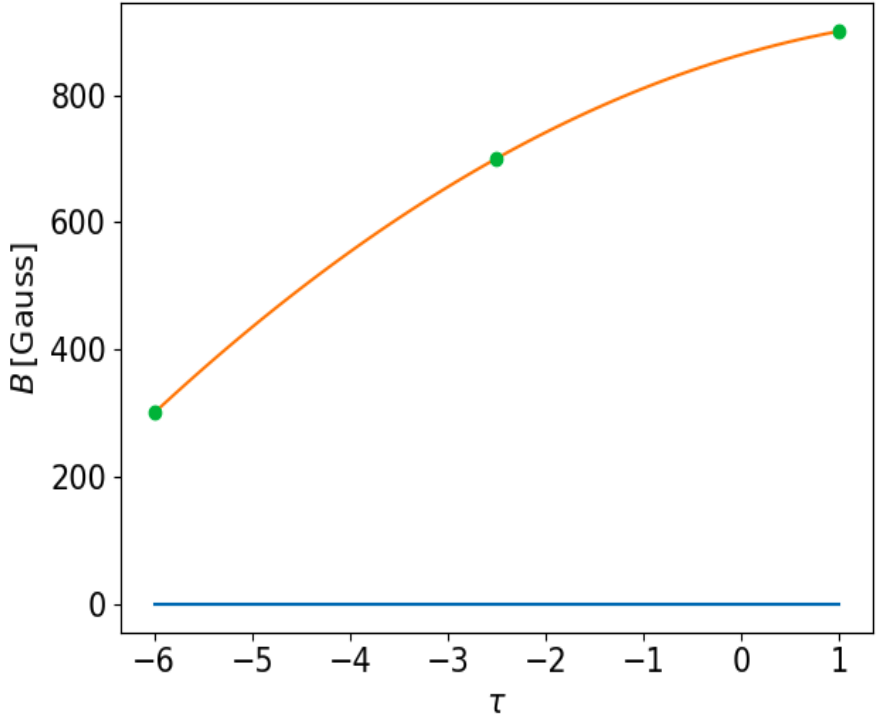
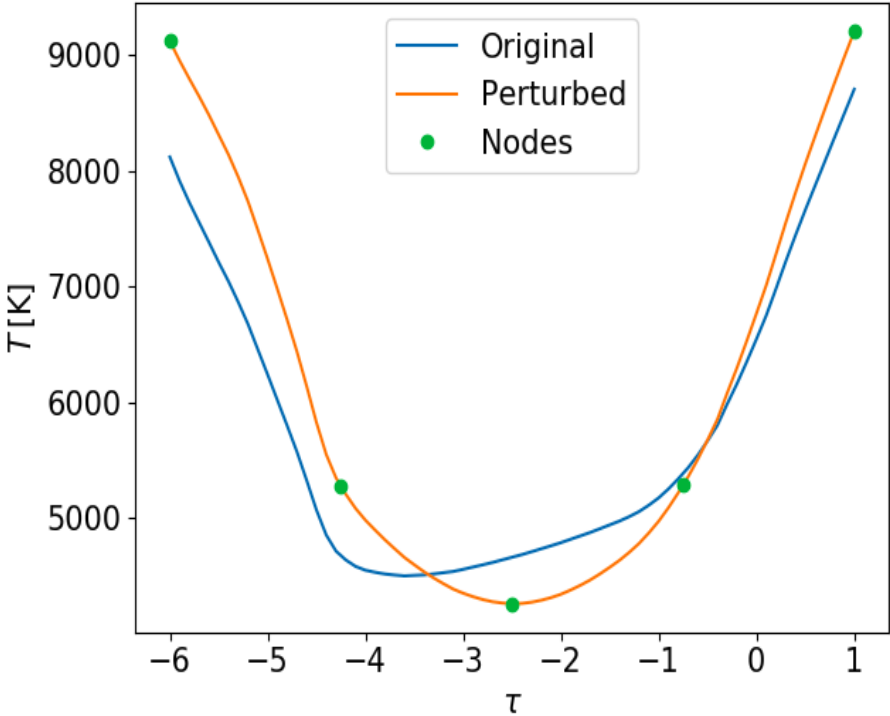
- Well, first of all, we can discretize our atmosphere as much as we want
- Does that mean all these values can be inferred? **No**
- Some points are simply not important.
- Some points are degenerate with other points.
- This means that inversion process is, actually, underdetermined.
- One way to think about this is that dimensionality of our data is low.

Our data is low dimensional: First 9 PCA components of our dataset



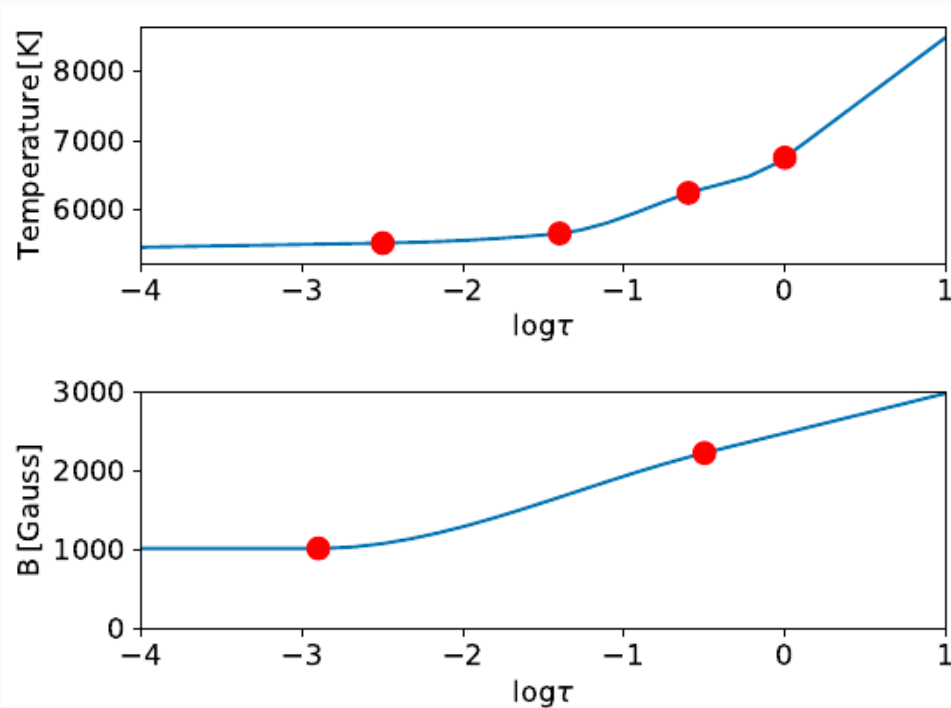
And that is why we use nodes

- You usually “reduce” number of parameters by including so called nodes. Still, the positions of the nodes / their number, come from some undefined scientific processes (personal preference, experience, etc.)

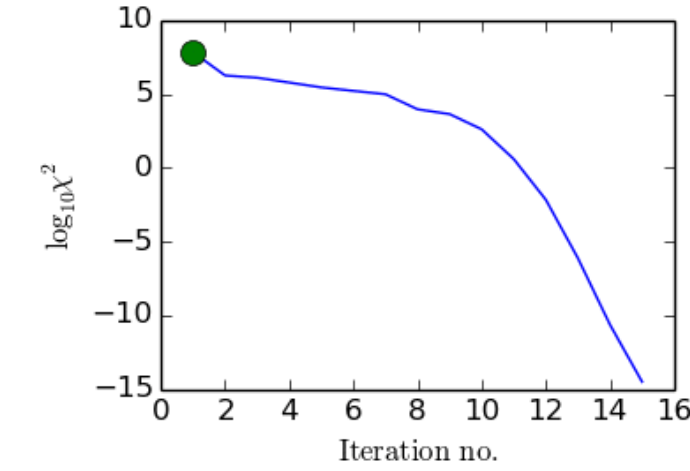
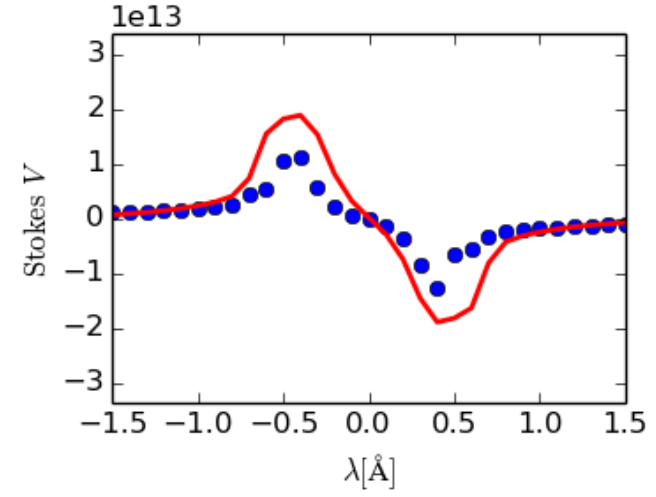
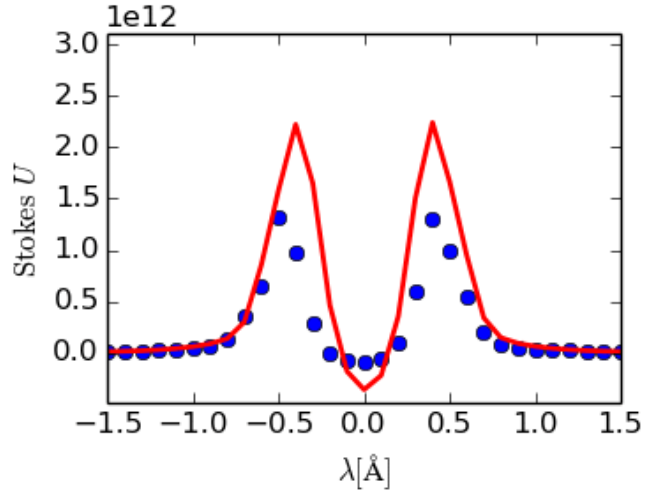
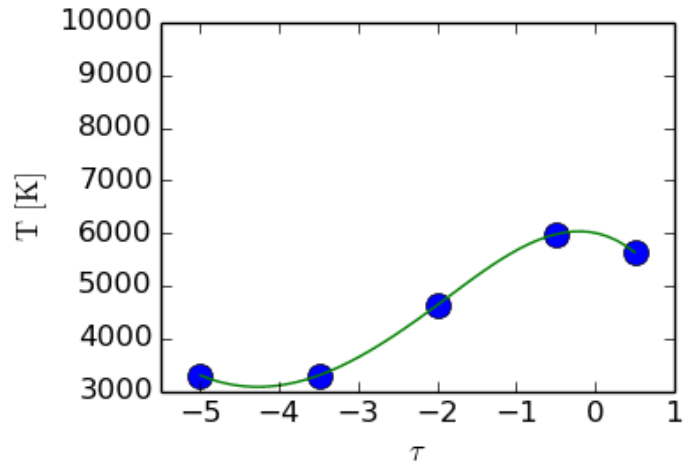
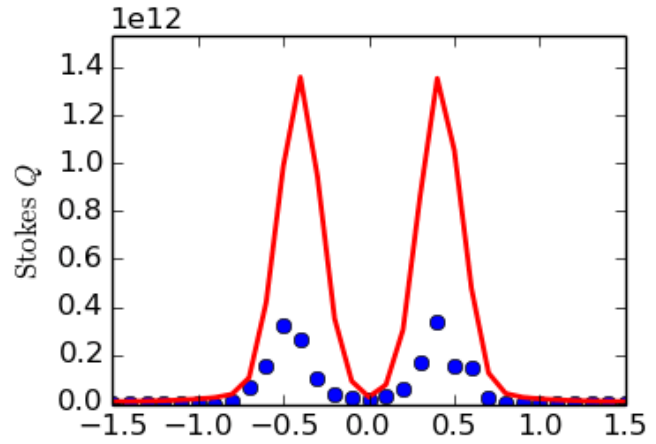
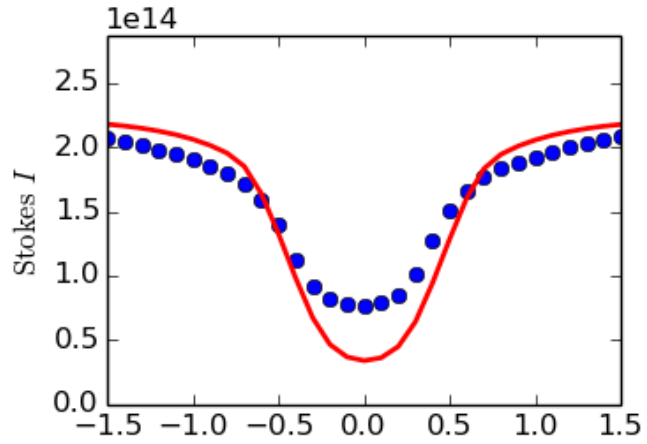


Instead of representing the full depth dependence of the atmosphere, we restrict it to few points. We map \mathbf{p} to α , where dimension of α is much smaller. Obviously, we need new response functions:

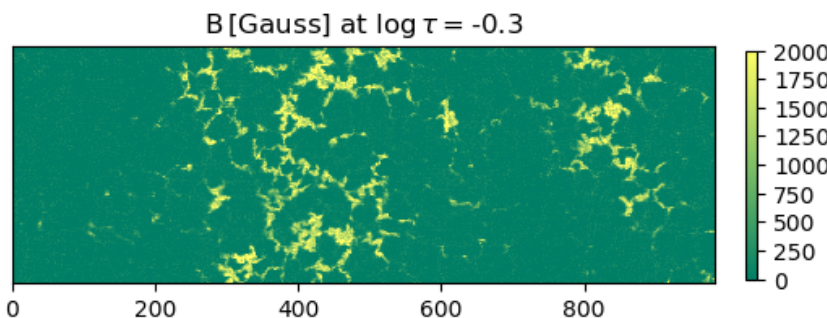
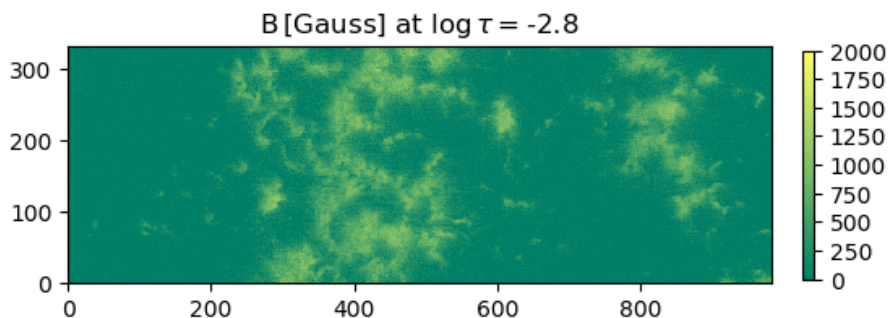
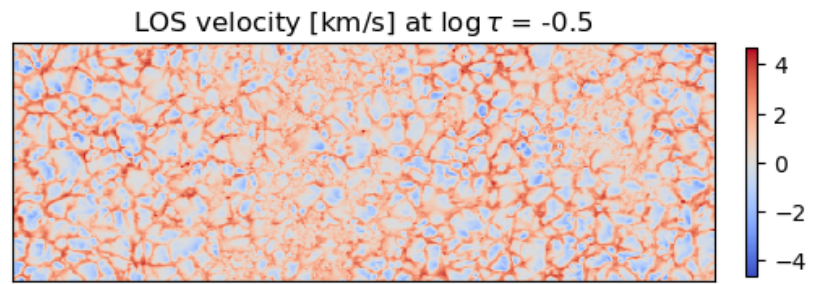
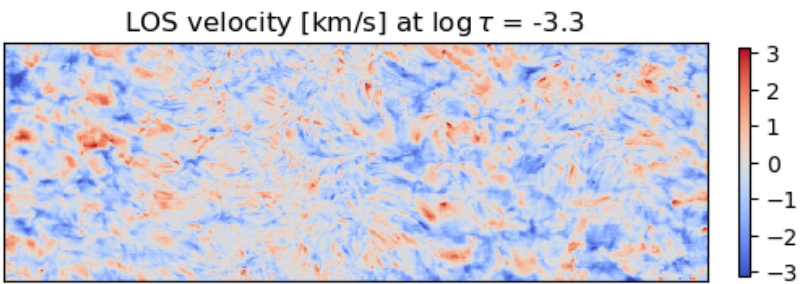
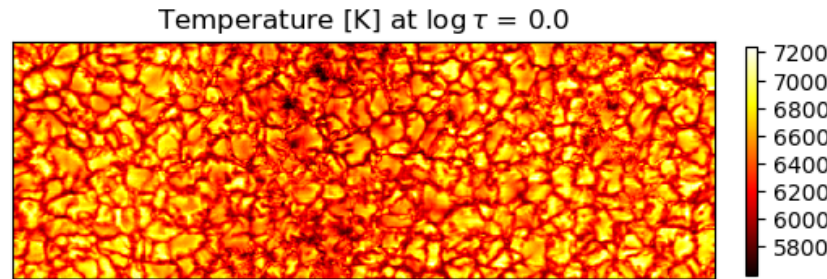
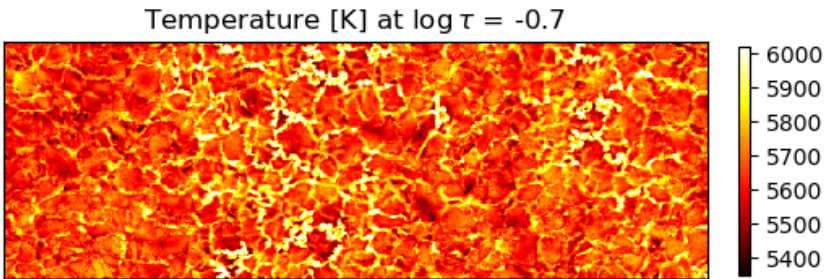
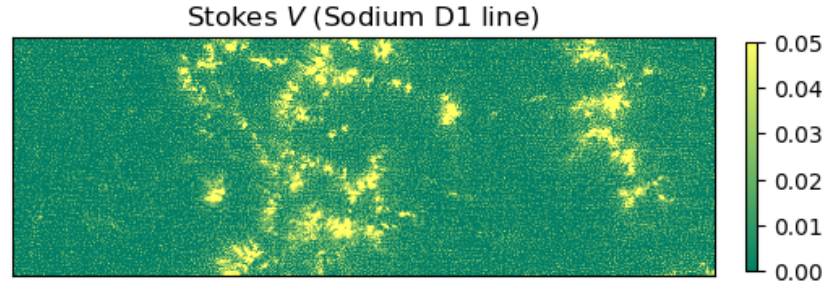
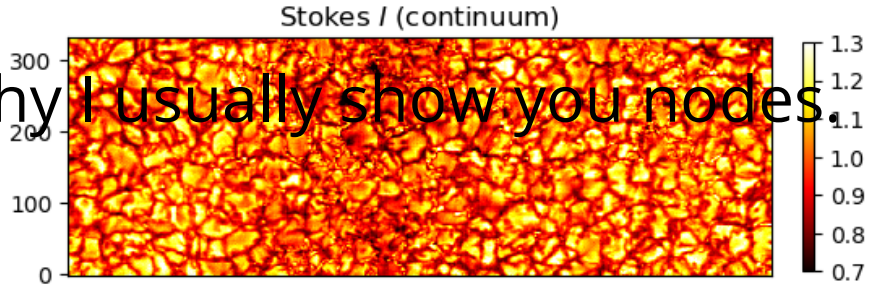
$$\frac{\partial l_i}{\partial \alpha_j} = \sum_{j'} \frac{\partial l_i}{\partial p_{j'}} \frac{\partial p_{j'}}{\partial \alpha_j}$$



So now we can understand what happens here



That is why I usually show you nodes.



What does this mean for you

- If you want to use DKIST, you will have to deal with inversions in some way
- Maybe you will do them yourself (several options)
- Or we will do them for you (then you will get a *cube* of your parameter models)
- So you will get a **model atmosphere** for each of your **observed spectra**
- It is good to know what came into that model atmosphere
- Keep eye out for **DKIST workshops** and **HAO Spectropolarimetry school**